

**<sup>10</sup>BE, <sup>26</sup>AL, AND <sup>53</sup>MN IN MARTIAN METEORITES.** C. Schnabel<sup>1,\*</sup>, P. Ma<sup>1</sup>, G.F. Herzog<sup>1</sup>, T. Faestermann<sup>2</sup>, K. Knie<sup>2</sup>, and G. Korschinek<sup>2</sup> <sup>1</sup>Dept. Chemistry, Rutgers Univ., New Brunswick, NJ 08854-8087, <sup>2</sup>Fakultät für Physik, Technische Universität München, 85748 Garching, Germany, \*Present address: ETH Hoenggerberg, Inst. Particle Physics, CH 8093 Zurich, (schnabel@particle.phys.ethz.ch)

**Introduction:** Cosmic-ray exposure ages of martian meteorites help determine how many separate events brought meteorites from Mars to Earth. The activities,  $A$ , of cosmogenic radionuclides give cosmic-ray exposure (CRE) ages ( $T$ ) when we know 1) the terrestrial age,  $t$ , of the meteorite; 2) the production rate,  $P$ , of the nuclide; and 3) that the period of exposure did not last more than  $\sim 3$  half-lives of the nuclide measured. We then have  $T = -\ln(1 - Ae^{\lambda t}/P)/\lambda$ , where  $\lambda$  is the decay constant. The CRE ages of many martian meteorites [1,2] approach or exceed the bound of  $\sim 3$  half-lives for the radionuclides most commonly determined. With its longer half-life of 3.7 My, <sup>53</sup>Mn records irradiation times up to  $\sim 11$  My. Advances in accelerator mass spectrometry (AMS) now allow the measurement of <sup>53</sup>Mn/<sup>55</sup>Mn ratios as low as  $5 \times 10^{-14}$  [3], thus reducing the mass of sample needed for analysis and making a survey of <sup>53</sup>Mn in SNC meteorites feasible.

**Experimental Methods:** We analyzed 100-to-200-mg samples of six SNC meteorites (Table 1) and 10-to-30 mg samples of several control samples, namely, Dhurmsala (LL6) for <sup>26</sup>Al and <sup>10</sup>Be; and Allende (CV3), Bogou (IA), ALH 77250 (IA), and Grant (IIB) for <sup>53</sup>Mn.

After the addition of appropriate carriers stony material was dissolved in HF, HNO<sub>3</sub>, and HClO<sub>4</sub> and irons in dilute HNO<sub>3</sub>. To separate the elements of interest, we evaporated the solution, took up the residue in 1 M HCl, took an aliquot for chemical analysis, added a few drops of 3% H<sub>2</sub>O<sub>2</sub>, and evaporated the solution to dryness. We dissolved this residue in 10.2 M HCl, loaded the solution on an anion exchange column, and eluted two fractions: Al + Be + Ni and most Cr<sup>3+</sup> with 10.2 M HCl; and Mn<sup>2+</sup> and any remaining Cr<sup>3+</sup> with 7.1 M HCl. To minimize Cr<sup>3+</sup>, which interferes in subsequent AMS, we purified the Mn-bearing fraction by repeating the foregoing procedure. After evaporation, the residue from the Mn fraction was dissolved in 7 M HNO<sub>3</sub> (ultra pure) and Mn was precipitated as MnO<sub>2</sub> on addition of KClO<sub>3</sub>. The MnO<sub>2</sub> was redissolved and precipitated, and dried, first at 110°C and then at 250°C. Procedures for the separation of Be and Al followed [4].

We used AMS to measure <sup>53</sup>Mn/<sup>55</sup>Mn ratios at the Technische Universität München [3]. Initially, results were normalized to those for a laboratory standard in which the <sup>53</sup>Mn had been produced by a nuclear reaction and its concentration determined by an

inbeam measurement of the deexcitation gamma rays of <sup>53</sup>Mn [5]. A procedural blank gave <sup>53</sup>Mn/<sup>55</sup>Mn  $\leq 1 \times 10^{-12}$  (atom/atom).

The <sup>10</sup>Be/<sup>9</sup>Be and <sup>26</sup>Al/<sup>27</sup>Al ratios were measured at PRIME Lab of Purdue University. We also analyzed several sample aliquots for elemental Mn, Fe, and Ni by ICP-MS.

**Results:** Elemental Mn and/or Fe contents (mass %) follow: ALH 77005, --, 15.4±0.8; ALH 84001, 0.369, 14.4±0.7; EET 79001A, 0.414, 14.9±0.7; LEW 88516, 0.360, 14.8±0.5; QUE 94201, 0.354%, 13.8±0.9; Zagami, --, 17.0±0.9. The results agree within  $\sim 10\%$  with literature values [6] except for Fe in Zagami (17.0 vs. 14.1 [6]). We also measured the Ni and Fe contents (mass %) of Bogou, 7.01±0.35 and 89.6±4.5, and of Grant Bar I, 8.71±0.44 and 88.8±4.4.

Analyses of <sup>53</sup>Mn in control samples gave the following results: ALH 77250, 347±39, 395±26; Bogou (USNM 2245) 351±71, 412±41; Bogou (KN) 368±37, 364±30 all in [dpm/[kg (Fe+1/3Ni)]]; Allende [dpm/(kg Fe)], 260±54, 184±65; Grant [dpm/kg], 367±70, 431±43. Literature values in the same units [see 6] are as follows: ALH 77250, 565±22; Allende (3529), 330±31; Bogou, 471±20; Grant, 373±10 (bar I). Except in Grant, our <sup>53</sup>Mn activities are systematically lower than published results by a factor of 1.33±0.12 ( $2\sigma$ ; weighted average; 8 measurements). A review of results for an internal laboratory standard for AMS, Sikhote-Alin (IIB), gave a similar factor of 1.27±0.13. We therefore apply a provisional factor of 1.33 to normalize all <sup>53</sup>Mn activities. Recalculated values for the meteorites above are as follows (normalization uncertainty not included): ALH 77250, 493±34; Bogou (USNM 2245), 507±55; Bogou (KN), 487±40; Allende, 295±72; and Grant, 531±57. Two separate Dhurmsala samples each contained 21.2±0.3 dpm <sup>10</sup>Be/kg and 67.6±3.3 and 73.7±2.2 dpm <sup>26</sup>Al/kg. Results for <sup>53</sup>Mn, <sup>10</sup>Be, and <sup>26</sup>Al in SNC meteorites appear in Table 1. The <sup>26</sup>Al and <sup>10</sup>Be activities agree well with published results [8-10].

**Discussion:** Table 1 shows published <sup>3</sup>He and <sup>21</sup>Ne CRE ages,  $T_3$  and  $T_{21}$  [1], and CRE ages calculated from measured radionuclide activities corrected for terrestrial age [1]. We omit  $T_{38}$  from the compilation because these results tend systematically to be about 15% lower than CRE ages based on <sup>3</sup>He and <sup>21</sup>Ne. Production rates are

**Table 1:**  $^{10}\text{Be}$ ,  $^{26}\text{Al}$ , (dpm/[kg meteorite]) and  $^{53}\text{Mn}$  (dpm/[kg (Fe)] activities and production rates (atom/min/kg or atom/min/[kg Fe]), and CRE ages (T[My]).

Meteorite	EET 79001A	QUE 94201	Zagami	ALH 77005	LEW 88516	ALH 84001
ID	522	48		21	38	339
$^{26}\text{Al}$	33.8±3.7	63.4±6.9	97.7±8.5	47.1±2.4	81.8±5.2	70.2±3.0
$^{10}\text{Be}$	4.99±0.07	11.9±0.2	14.6±0.2 <sup>[8]</sup>	16.2±0.8 <sup>[8]</sup>	16.2±0.2	21.3±0.3
$^{53}\text{Mn}$	30±5	162±11	221±16	170±12	279±20	379±33
$T_{\text{terr}}^{[1]}$	0.2	0.3		0.2		
D, R (cm)	15,15	25,25	25,25	10,10	25,25	25,6
$P_{26}$	72	100	93	49	77	70
$P_{10}$	20	21	21	19	22	21
$P_{53}$	346	451	451	284	451	376
$T_{26}$	0.86±0.21	1.9±0.9	--	--	--	--
$T_{10}$	0.70±0.06	2.2±0.3	2.5±0.2	4.2±1.3	3.0±0.3	--
$T_{53}$	0.51±0.10	2.6±0.3	3.6±0.4	5.2±0.8	5.1±0.8	--
$T_3^{[1]}$	0.61	2.1	2.9	3.8-4.5	4.3	15
$T_{21}^{[1]}$	0.65	3.2	3.0	3.2-4.9	4.3	14.2

calculated from elemental compositions [6] and elemental production rates [11].

*EET 79001* – CRE ages from noble gases agree at about 0.6 My. Our  $^{53}\text{Mn}$  activity, 30±5, is lower than reported values of 60-65 dpm/[kg Fe] (see [7]). Estimates of the preatmospheric radius, R, of EET 79001 range from 10 to 15 cm. Assuming that our sample came from a 15 cm body [12] at a depth, D=15 cm, large enough to preclude SCR production, we obtain CRE ages in fair agreement with those based on noble gases. Smaller depth for fixed radius would increase  $T_{26}$ ,  $T_{10}$ , and  $T_{53}$ .

*QUE 94201* – Nishiizumi and Caffee [8] report 1)  $^{10}\text{Be}$  and  $^{26}\text{Al}$  activities similar to ours; and 2) little evidence for the effects of SCR. We assume a sample depth of 25 cm and a meteoroid radius of 25 cm because these choices correspond to normal production rates; smaller depths in bodies with radii of 30 cm to 40 cm give similar results. The various CRE ages are in reasonable agreement.

*Zagami* – Noble gas CRE ages are close to 3.0 My. Again assuming standard production rates, the  $^{10}\text{Be}$  age based on the analysis of [9] gives a similar result;  $T_{53}$  is higher.

*ALH 77005* – Solar cosmic ray effects and track data combined with an assumed CRE age of 2.5 My originally suggested a very small preatmospheric radius of 4-6 cm [13] for ALH 77005. Absent SCR effects, GCR production rates in such small objects are depressed:  $P_{10}<17$  ( $^{10}\text{Be}$  ~saturated) and  $P_{53}<200$  atom/min/[kg Fe] [11]. The  $^{53}\text{Mn}$  CRE calculated with  $P_{53}=200$ , ~12 My, is much larger than values

based on noble gases, 3-5 My. We therefore assume a somewhat larger but still small body with R=10 cm. As the low  $^{26}\text{Al}$  activity in our sample indicates little or no SCR production, we further assume an interior location. The resulting CRE ages are in the range estimated from the noble gases. A larger radius would raise production rates and lower the CRE ages.

*LEW 88516* – Lacking information about size, we adopt production rates for

the center of a 25-cm object.  $T_{53}$  agrees with the noble gas CRE ages;  $T_{10}$  is lower.

*ALH 84001* - Noble gas CRE ages of ~14 My imply a 4 $\pi$  irradiation long enough to have saturated  $^{10}\text{Be}$ ,  $^{26}\text{Al}$  and  $^{53}\text{Mn}$ . We adopt a radius of 25 cm, which is consistent with the dimensions of the recovered mass, and a depth of 6 cm based on a comparison of  $^{10}\text{Be}$  contents measured by us and by [9]. These choices lead to production rates that agree with the measured activities.

**Conclusion:** CRE ages of SNC meteorites based on  $^{53}\text{Mn}$  activities generally agree with CRE ages based on other cosmogenic nuclides. Shergottite basalts QUE 94201 and Zagami came from a more recent ejection event than did lherzolites ALH 77005 and LEW 88516.

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